

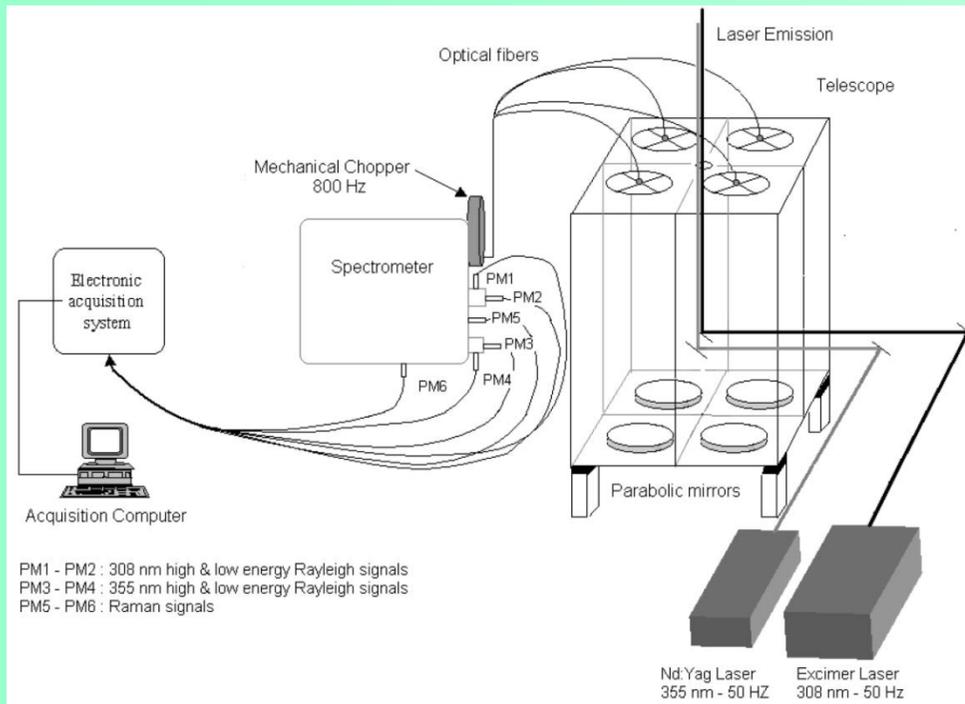


The effect of change of BP to DBM ozone absorption cross-sections on lidar measurements

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Principle of lidar ozone measurement

DIAL Method : Differential Absorption Lidar



Common wavelengths pairs used:

stratospheric systems: 308, 351-355 nm

Tropospheric: 266, 289, 299, 316 nm

- ✓ Emission of 2 laser beams in the UV range ($\lambda_{on}, \lambda_{off}$)
- ✓ Different ozone absorption cross-section
- ✓ Pulsed laser sources: range resolved measurement
- ✓ Large dynamic of the lidar signals: several acquisition channels
- ✓ N₂ Raman wavelengths: volcanic aerosols
- ✓ Self calibrated measurement

Retrieval of ozone number density: Rayleigh

DIAL Method : Differential Absorption Lidar

$$n_{O_3}(z) = -\frac{1}{2 \underbrace{\Delta\sigma_{O_3}(z)}_{\substack{\text{Differential ozone} \\ \text{absorption cross-section} \\ \sigma_{O_3}(\lambda_{on}, z) - \sigma_{O_3}(\lambda_{off}, z)}}} \frac{d}{dz} \text{Ln} \left(\frac{\underbrace{S(\lambda_{on}, z)}_{\text{lidar signal}} - \underbrace{S_b(\lambda_{on}, z)}_{\text{background}}}{\underbrace{S(\lambda_{off}, z)}_{\text{lidar signal}} - \underbrace{S_b(\lambda_{off}, z)}_{\text{background}}} \right) + \underbrace{\delta n_{O_3}(z)}_{\text{correction term}}$$

Laser wavelengths chosen so that the correction term is less than 10% of main term

$$\delta n_{O_3}(z) = \frac{1}{\Delta\sigma_{O_3}(z)} \left[\frac{1}{2} \frac{d}{dz} \text{Ln} \left(\frac{\underbrace{\beta(\lambda_{on}, z)}_{\substack{\text{backscatter} \\ \text{Rayleigh \& Mie}}}}{\underbrace{\beta(\lambda_{off}, z)}_{\substack{\text{backscatter} \\ \text{Rayleigh \& Mie}}}} \right) - \underbrace{\Delta\alpha(z)}_{\substack{\text{extinction} \\ \text{Rayleigh \& Mie}}} - \sum_i \underbrace{\Delta\sigma_i n_i(z)}_{\substack{\text{Extinction by} \\ \text{other species}}} \right]$$

Retrieval of ozone number density: Raman

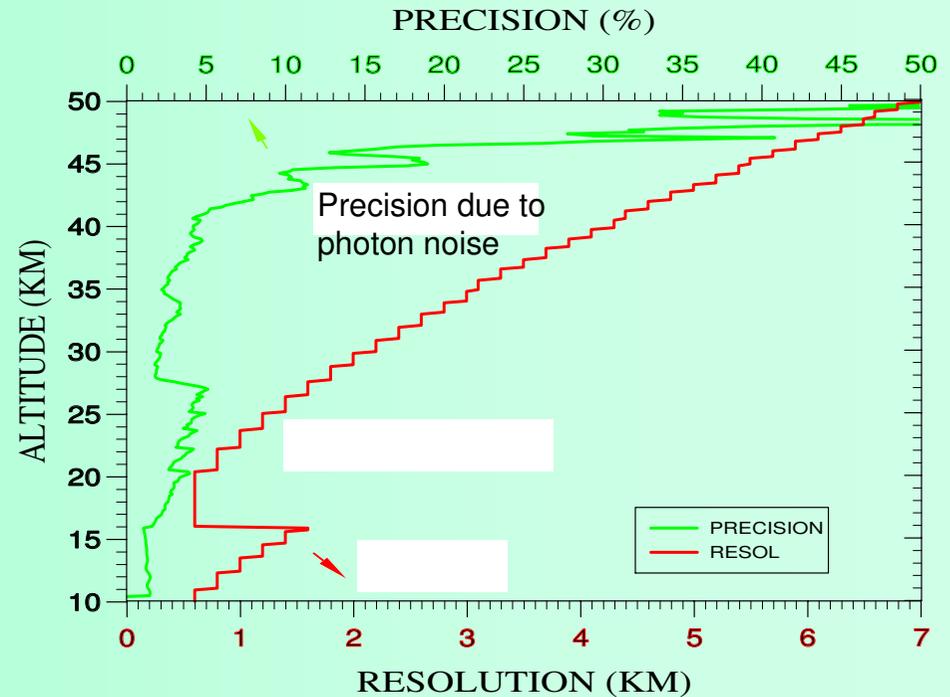
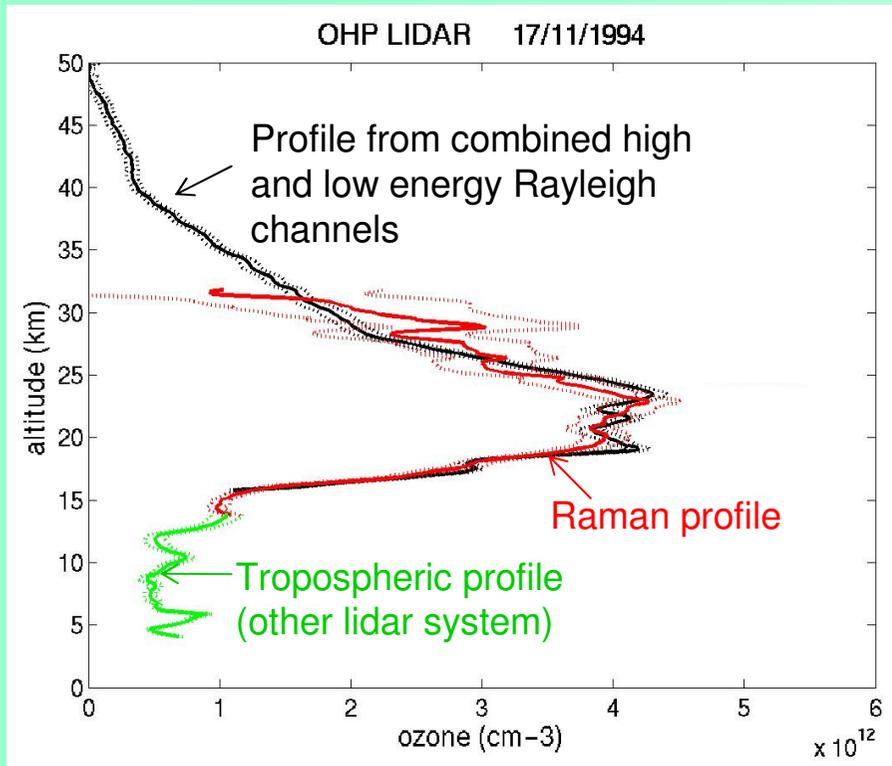
Use of Raman signals in the presence of volcanic aerosols

$$n_{O_3}^R(z) = - \frac{1}{\underbrace{\Delta\sigma_{O_3}^R(z)}_{\text{Differential ozone absorption cross-section}}} \frac{d}{dz} \text{Ln} \left(\frac{\underbrace{S(\lambda_{on}^R, z)}_{\text{lidar signal}} - \underbrace{S_b(\lambda_{on}^R, z)}_{\text{background}}}{\underbrace{S(\lambda_{off}^R, z)}_{\text{lidar signal}} - \underbrace{S_b(\lambda_{off}^R, z)}_{\text{background}}} \right) + \underbrace{\delta n_{O_3}^R(z)}_{\text{correction term}}$$

$$\Delta\sigma_{O_3}^R = \sigma_{O_3}(\lambda_{on}, z) - \sigma_{O_3}(\lambda_{off}, z) + \sigma_{O_3}(\lambda_{on}^R, z) - \sigma_{O_3}(\lambda_{off}^R, z)$$

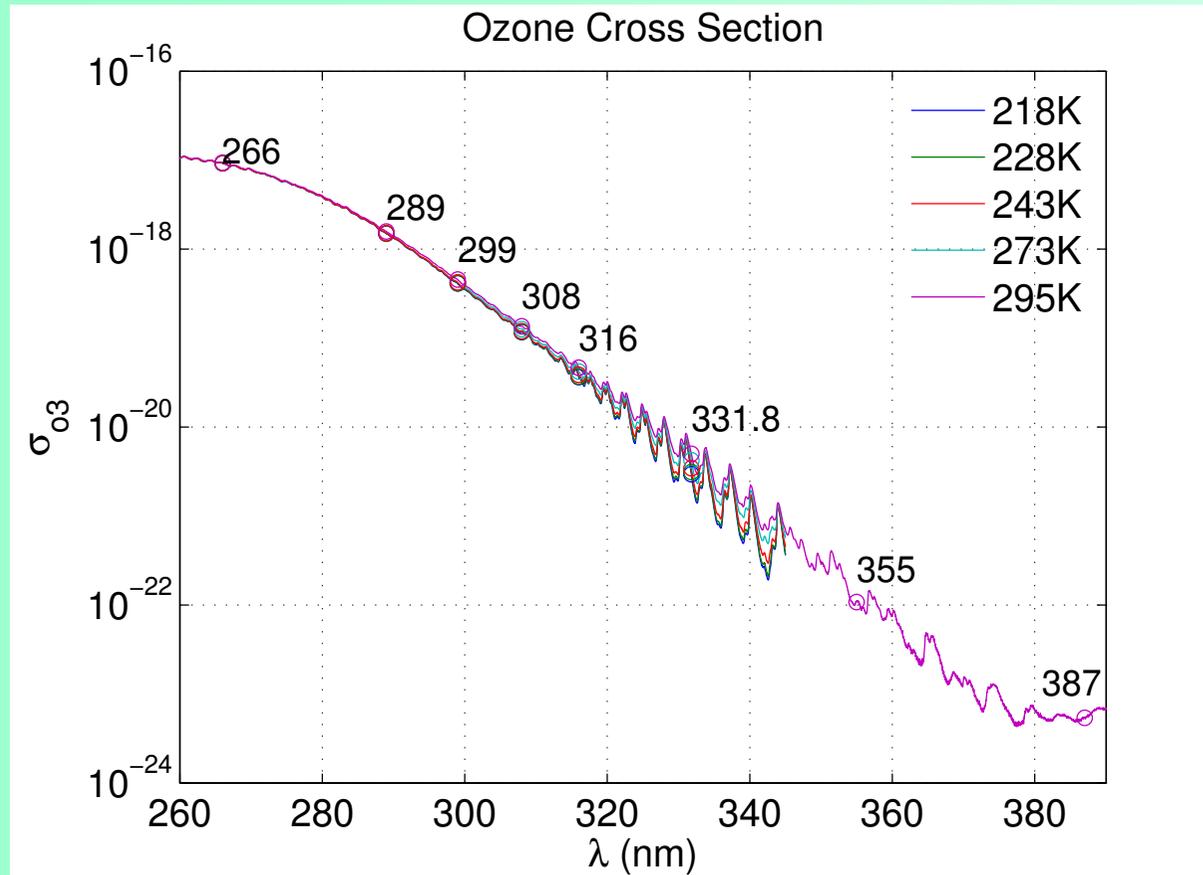
$$\delta n_{O_3}^R(z) = \frac{1}{\Delta\sigma_{O_3}^R(z)} \left[\underbrace{-\Delta\alpha^R(z)}_{\text{extinction Rayleigh \& Mie}} - \sum_i \underbrace{\Delta\sigma_i^R n_i(z)}_{\text{extinction by other species}} \right]$$

Example of DIAL ozone profile



- Ozone measurements performed during the night
- Temporal resolution 3 – 4 hours, depending on laser power and repetition rate
- Require clear skies

Wavelengths used in lidar measurements



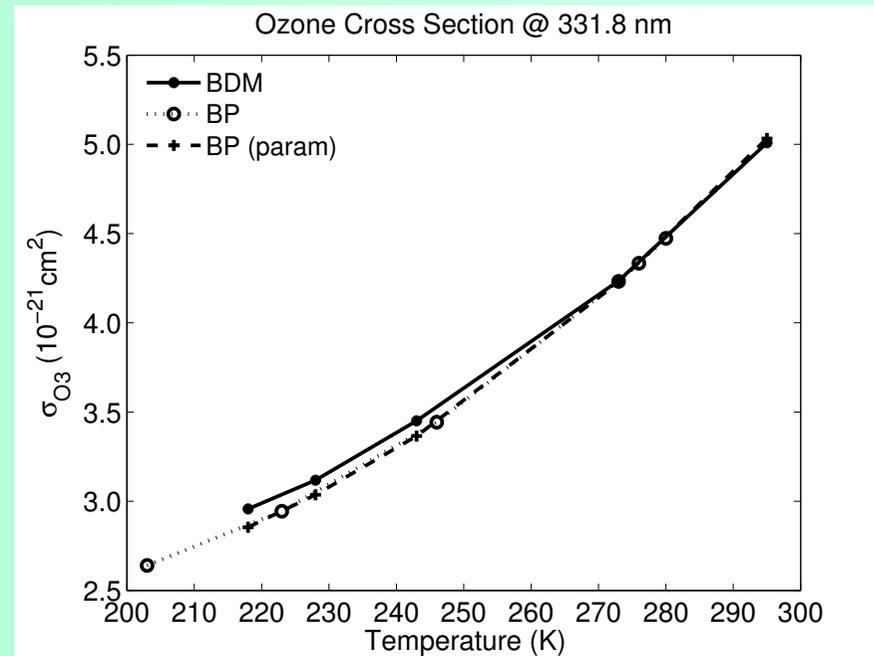
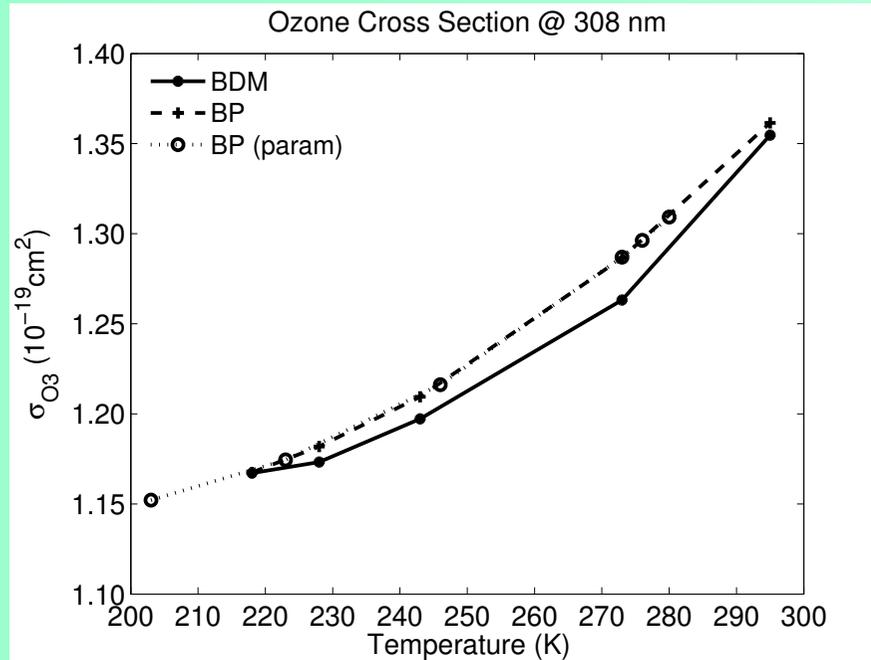
Spectral variation of ozone cross-section between 260 nm and 390 nm (Malicet et al., 1995; Brion et al., 1998).

Wavelengths used for ozone measurements in the troposphere and stratosphere.

Variation of ozone X-section with temperature

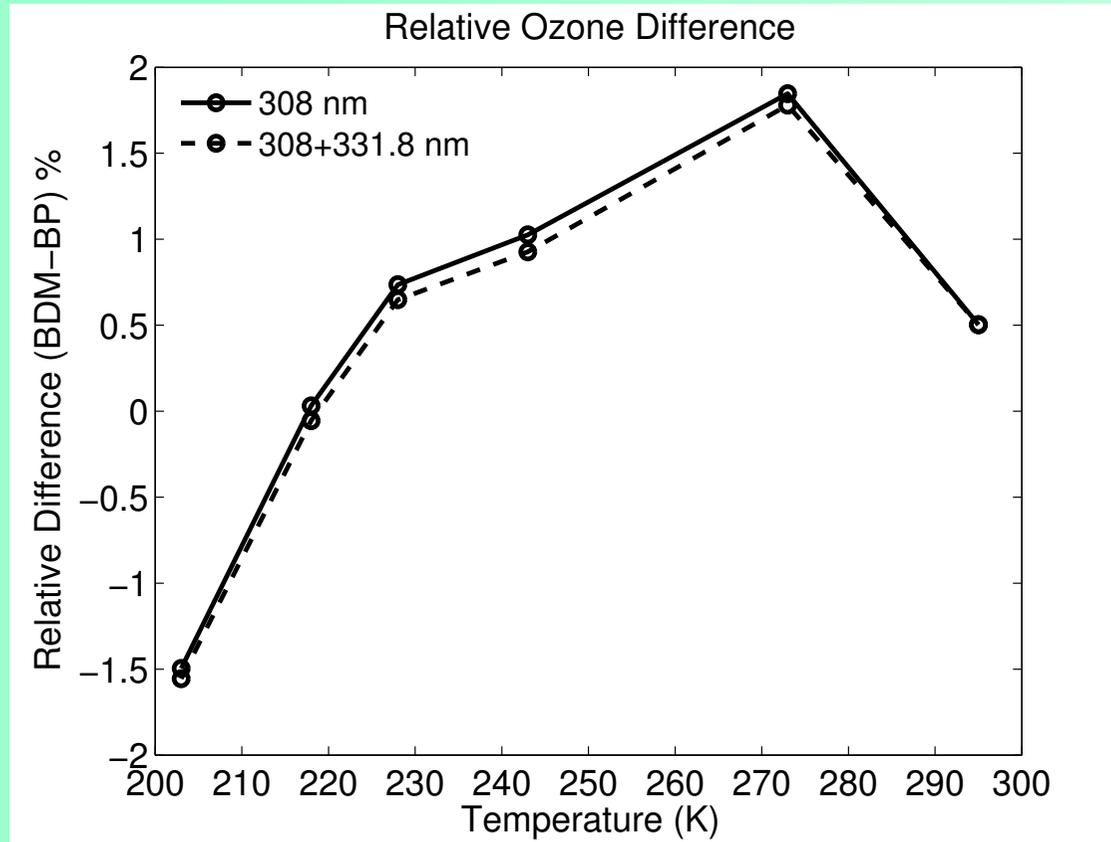
Rayleigh absorbed wavelength: 308 nm

Raman absorbed wavelength: 331.8 nm



Ozone cross-sections as a function of temperature for Malicet et al. (1995, BDM) and Bass and Paur (1984, BP) data sets.

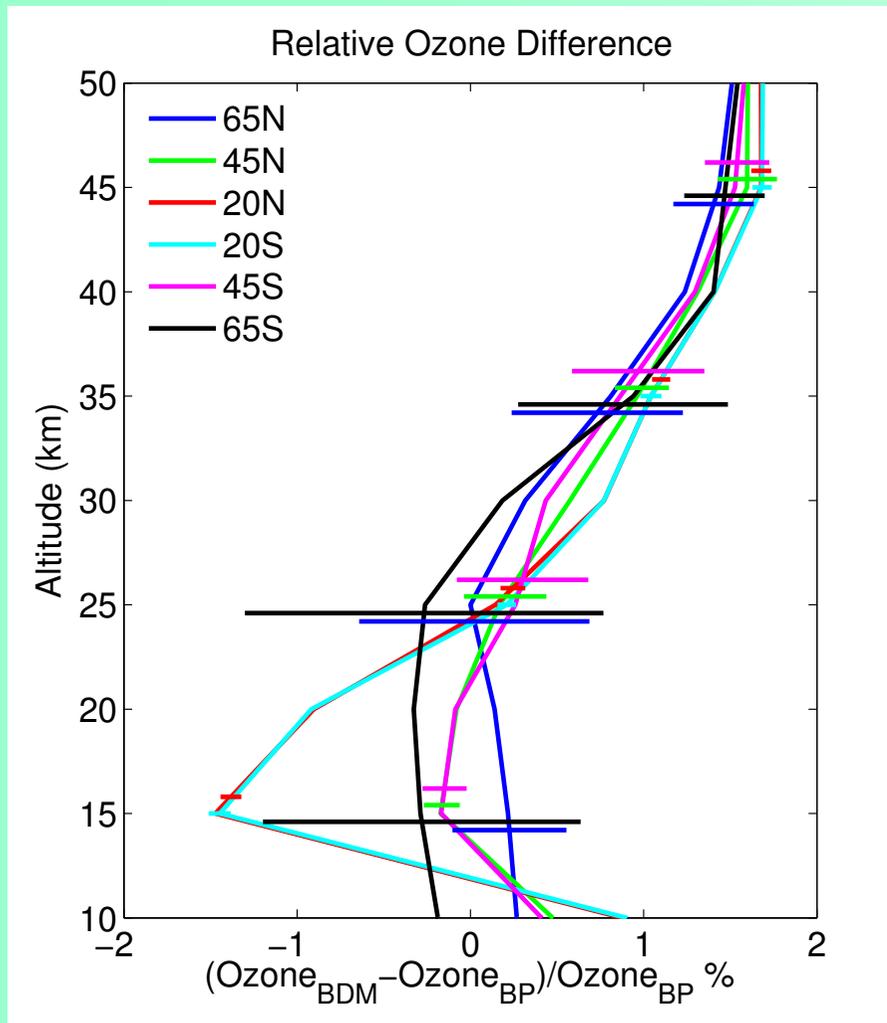
Effect of a change of X-sections on ozone



$$\Delta n_{O_3}(z) = \frac{\Delta \sigma_{O_3}^{BP}(z)}{\Delta \sigma_{O_3}^{BDM}(z)} - 1$$

Relative difference in ozone (BDM-BP)

Ozone change as a function of latitude



Relative difference in ozone (BDM-BP) using the CIRA model, for various latitude bands

Annual average

$$\overline{\Delta n_{O_3}(z)} = \sum_{m=1}^{12} \frac{n_{O_3}^{BDM}(z, m) - n_{O_3}^{BP}(z, m)}{n_{O_3}^{BP}(z, m)}$$

Conclusions

- The difference on ozone in both the classical DIAL retrieval based on elastic scattering and the retrieval based on Raman scattering is **below 1.5 %** in absolute value **from 10 to 30 km**.
- **Above 30 km**, the difference, estimated for the classical DIAL retrieval only, **are maximum around 45 km**. Largest differences are found **in the tropics** and reach about **1.8 %**.
- Correct evaluation of ozone cross-section temperature dependence is important for ozone trends evaluation from lidar measurements, due to temperature trends in the stratosphere.
- At present, most lidar groups within NDACC use Bass & Paur ozone cross-sections but DIAL ozone measurements can easily be re-computed from archived raw data in case of change in recommended ozone cross-section.

Initiative from the NDACC lidar group

Proposal in Response to the 2010 Call for the Formation of an ISSI International Team in Space Science on:

Critical Assessment and Standardized Reporting of Vertical Filtering and Error Propagation in the Data Processing Algorithms of the NDACC Lidars

Submitted by T. Leblanc, JPL

Ozone cross-section issue included.

