



# Sensitivity of stratospheric lidar ozone measurements to a change in ozone cross sections

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# Retrieval of ozone number density

## DIAL Method : Differential Absorption Lidar

$$n_{O_3}(z) = -\frac{1}{2 \cdot \Delta\sigma_{O_3}(z)} \frac{d}{dz} \text{Ln} \left( \frac{S(\lambda_{on}, z) - S_b(\lambda_{on}, z)}{S(\lambda_{off}, z) - S_b(\lambda_{off}, z)} \right) + \delta n_{O_3}(z)$$

Differential ozone
lidar signal
background

absorption cross-section

$$\sigma_{O_3}(\lambda_{on}, z) - \sigma_{O_3}(\lambda_{off}, z)$$

Laser wavelengths chosen so that the correction term is less than 10% of main term

$$\delta n_{O_3}(z) = \frac{1}{\Delta\sigma_{O_3}(z)} \left[ \frac{1}{2} \frac{d}{dz} \text{Ln} \left( \frac{\beta(\lambda_{on}, z)}{\beta(\lambda_{off}, z)} \right) - \Delta\alpha(z) - \sum_i \Delta\sigma_i n_i(z) \right]$$

correction term
Backscatter
Extinction
Extinction by

Rayleigh & Mie
other species

## Use of Raman signals in the presence of volcanic aerosols

$$n_{O_3}(z) = -\frac{1}{\Delta\sigma_{O_3}^R(z)} \frac{d}{dz} \text{Ln} \left( \frac{S(\lambda_{on}^R, z) - S_b(\lambda_{on}^R, z)}{S(\lambda_{off}^R, z) - S_b(\lambda_{off}^R, z)} \right) + \delta n_{O_3}^R(z)$$

Differential ozone
lidar signal
background

absorption cross-section

$$\sigma_{O_3}(\lambda_{on}, z) - \sigma_{O_3}(\lambda_{off}, z) + (\sigma_{O_3}(\lambda_{on}^R, z) - \sigma_{O_3}(\lambda_{off}^R, z))$$

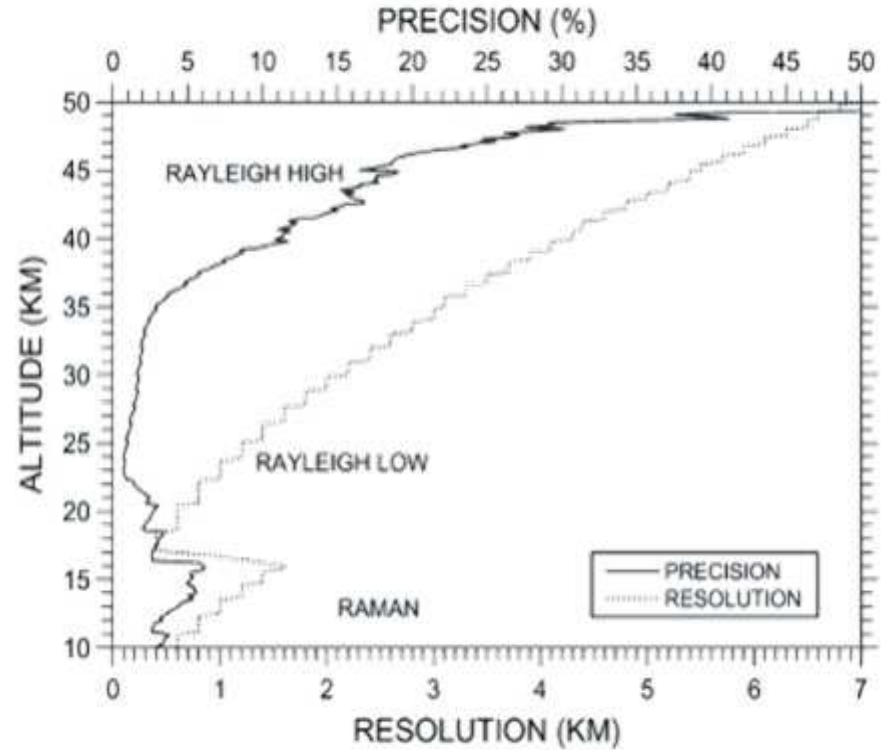
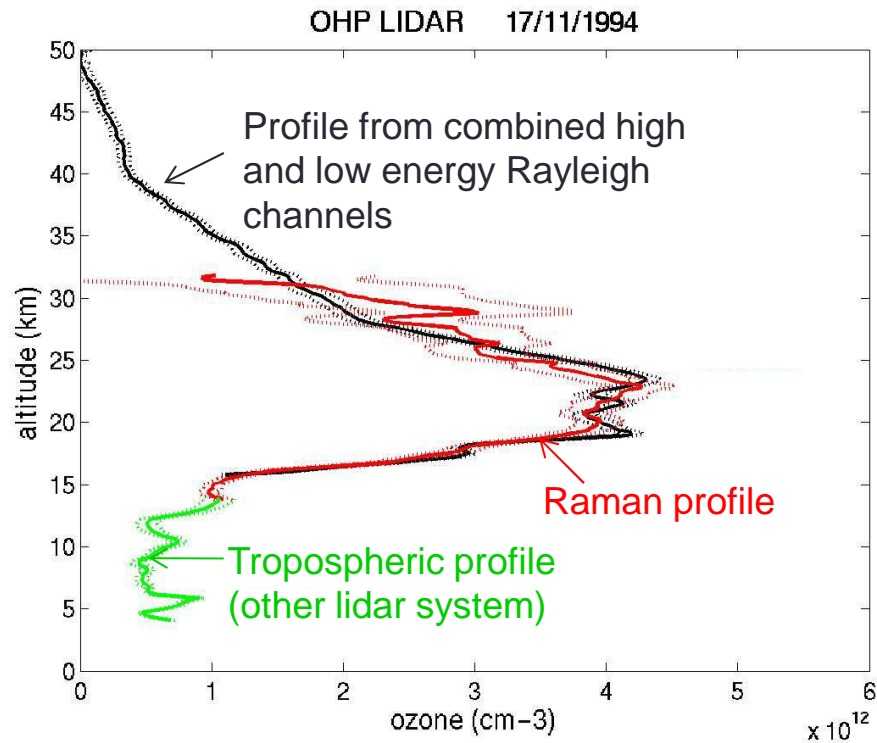
$$\delta n_{O_3}^R(z) = \frac{1}{\Delta\sigma_{O_3}^R(z)} \left[ -\Delta\alpha^R(z) - \sum_i \Delta\sigma_i^R n_i(z) \right]$$

correction term
Extinction
Extinction by

Rayleigh & Mie
other species

➡ A change in ozone cross-section data will affect the retrieved ozone number density

# Example of DIAL ozone profile



Ozone measurements performed during the night

Temporal resolution 3 – 4 hours, depending on laser power and repetition rate

Require clear skies

# Wavelengths used in lidar measurements

- ✓ DIAL ozone retrieval requires determination of ozone cross-sections in the range 266 nm – 387 nm, with accurate temperature dependence.
- ✓ For ozone observations, the selected pair of laser wavelengths differs according to the altitude range of the measurements
- ✓ For tropospheric measurements, the wavelength pair is chosen between 266 nm and 320 nm.
- ✓ For stratospheric measurements:
  - Most instrumental set-ups are based on XeCl excimer laser (XeCl) sources that emit several radiations in the range 307.9 -308.2 nm (Rayleigh channel)
  - In the presence of volcanic aerosols, the range of wavelength in the Raman lidar signal is 331.7-332.1 nm

# Ozone X-section used in this study

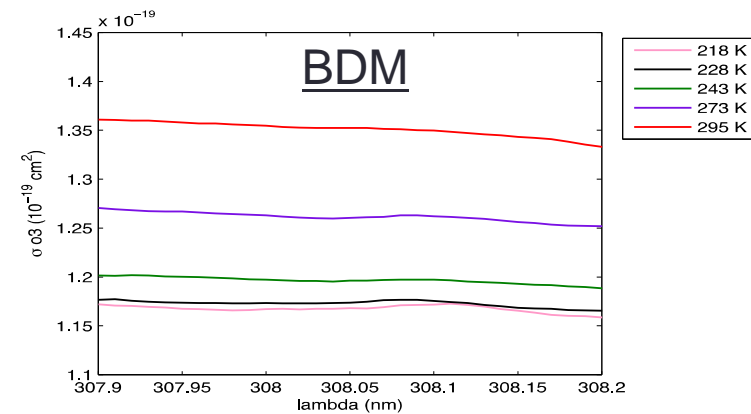
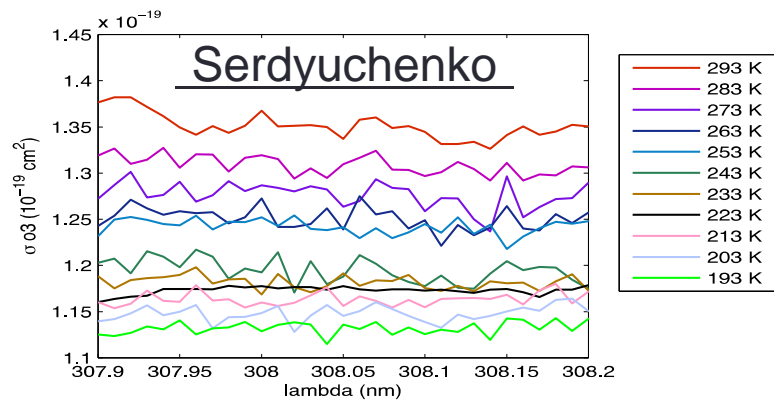
	Temperature ( K ) (number of T level )	Spectral Resolution(nm)
Bass and Paur (1984)	203 – 280 (6)	0.05
Brion Daumont and Malicet (1995)	218- 295 (5)	0.01
Serdyuchenko (2012)	193- 293 (11)	0.01

Serdyuchenko should be more appropriate for lidar measurements :

- Larger Temperature range
- Larger Temperature sampling

## Inconvenience of Serdyuchenko cross section:

O<sub>3</sub> cross section as a fonction of Lambda in lidar rayleigh absorbed wavelength

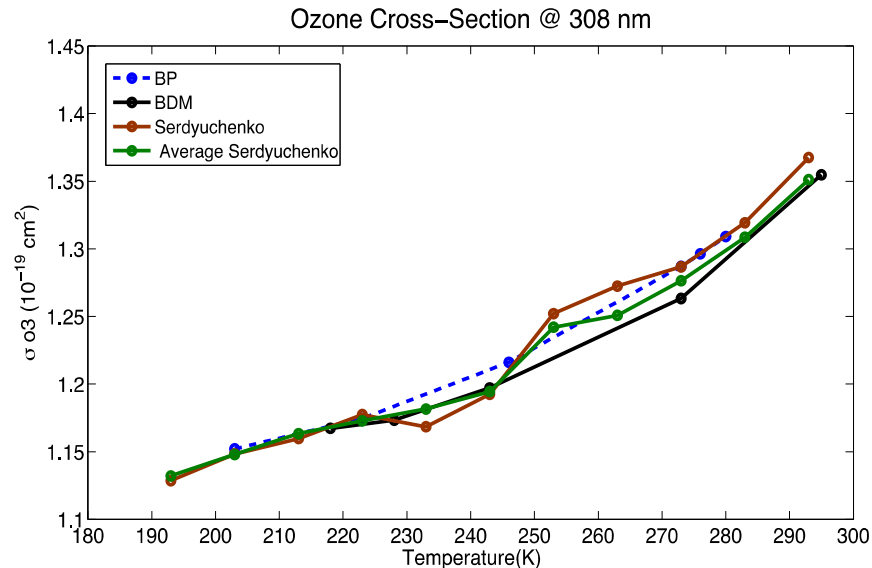


Serdyuchenko cross section noisier with respect to BDM

# Variation of ozone X-section with temperature

## Rayleigh absorbed wavelength: 308 nm

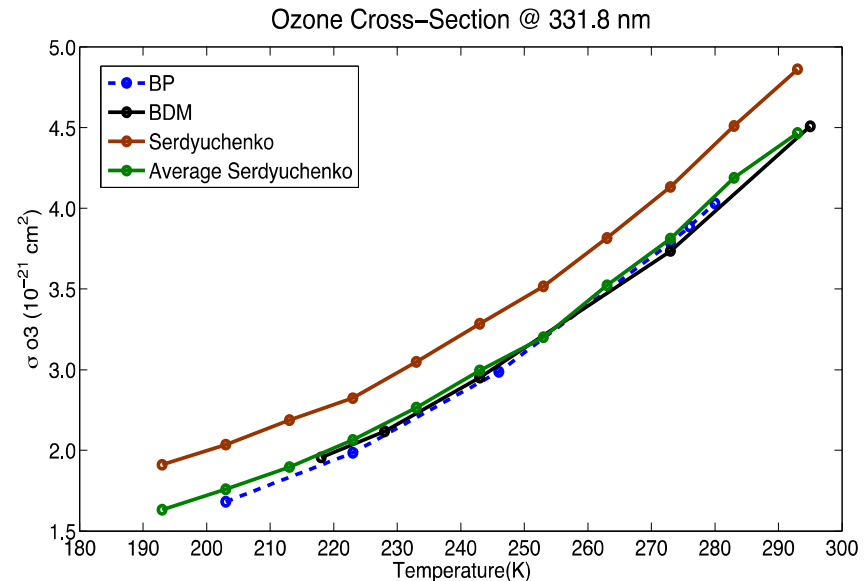
Average Serdyuchenko: mean (307.9 - 308.2 nm)



- Serdyuchenko X section Jump ( $0.1 \cdot 10^{-21} \text{ cm}^2$ ) at 242K
- Jump less seen on Average Serdyuchenko

## Raman absorbed wavelength: 331.8 nm

Average Serdyuchenko: mean (331.7 - 332.1 nm)



- Serdyuchenko ( at 331.8 nm) higher than BP constant bias  $\sim 0.4 \cdot 10^{-21} \text{ cm}^2$

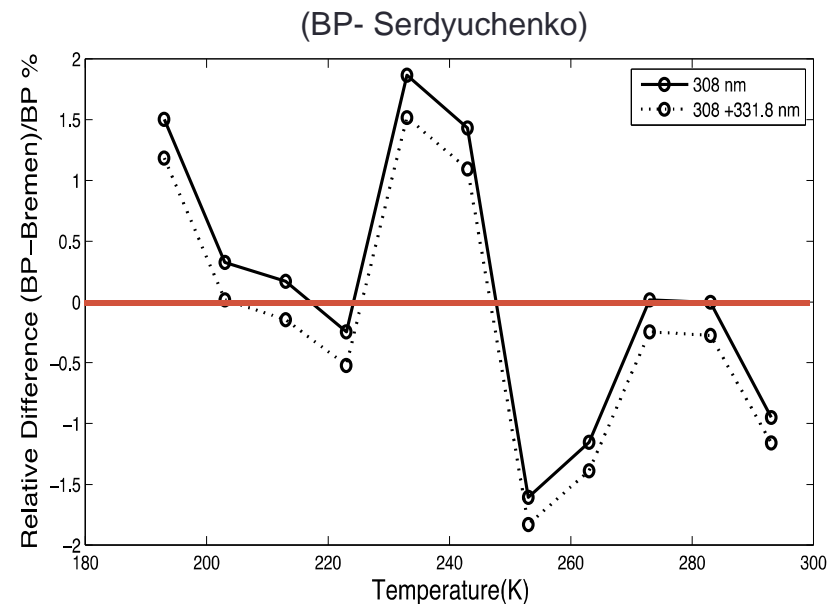
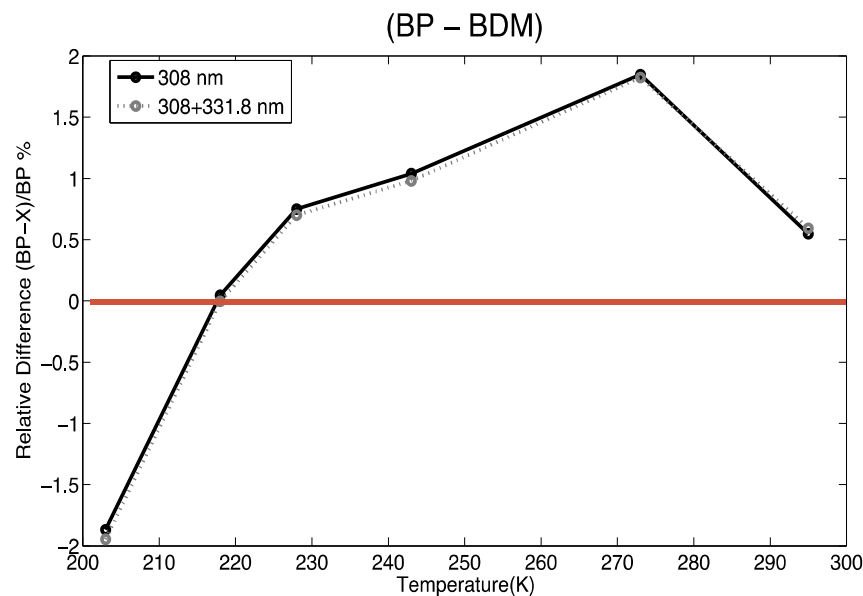


To compare each cross section an extrapolation of Temperature has to be made

	BP	BDM	Serdyuchenko	Average Serdyuchenko
Original range	203 – 280 K	218- 295 K	193- 293 K	193- 293 K
BP-X		203-295 K	193-293 K	193-293 K

# Effect of a change of X-sections on ozone

## Bass and Paur used as the reference

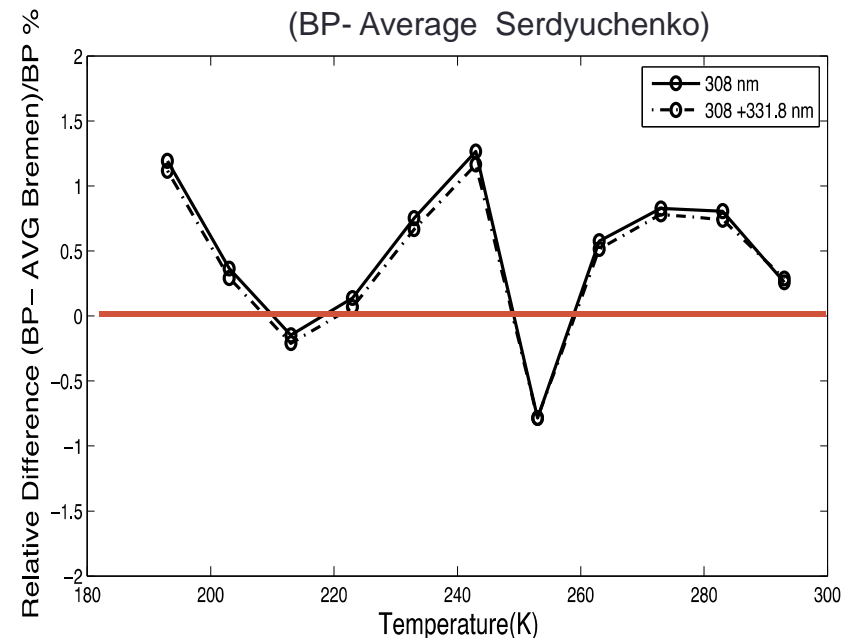


Impact on the X-sections more important on the rayleigh channel (308 nm)

At 308 nm

	BP-BDM	BP-Serdyuchenko	BP- Average Serdyuchenko
Min	-1.9% at 203 K	-1.7% at 253 K	0.7% at 203 K
Max	1.9% at 273K	1.9% at 233 K	1.2% at 233 K

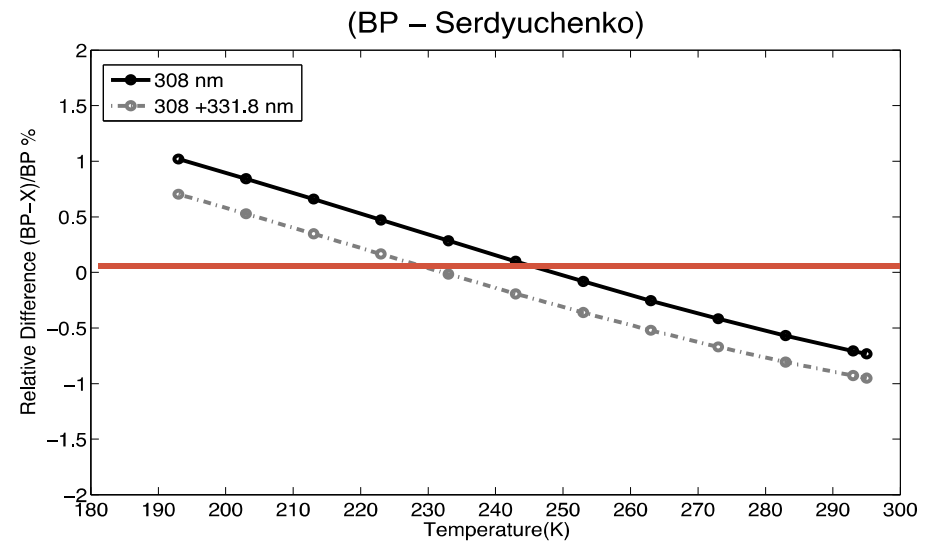
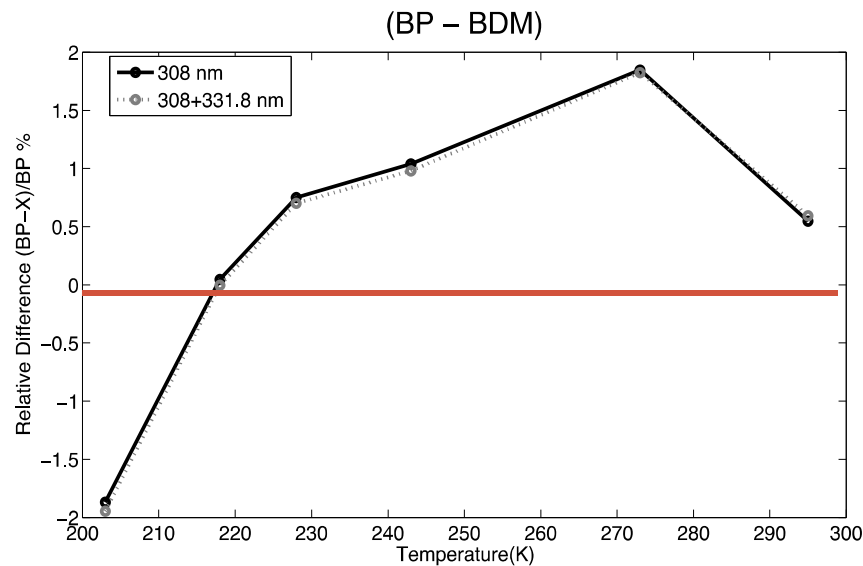
- Stronger amplitude of the difference with BDM
- Strong impact of the Serdyuchenko noise in the relative ozone difference



# Effect of a change of X-sections on ozone

## Bass and Paur used as the reference

➔ Reduction of the noise : An order 2 polynomial fit applied on Serdyuchenko sections

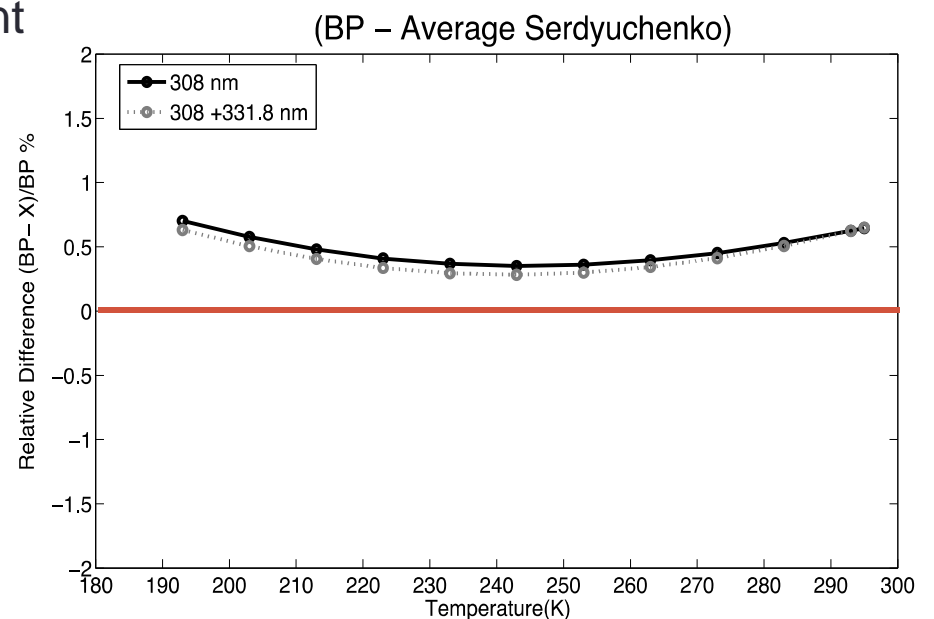


➔ Impact on the X-sections still more important on the rayleigh channel (308 nm)

At 308 nm

	BP-BDM	BP-Serdyuchenko	BP- Average Serdyuchenko
Min	-1.9% at 203 K	-0.7% at 295 K	0.3% at 243 K
Max	1.9% at 273K	1% at 193 K	0.7% at 193 K

- Stronger amplitude of the difference with BDM
- Reduction of the Serdyuchenko noise



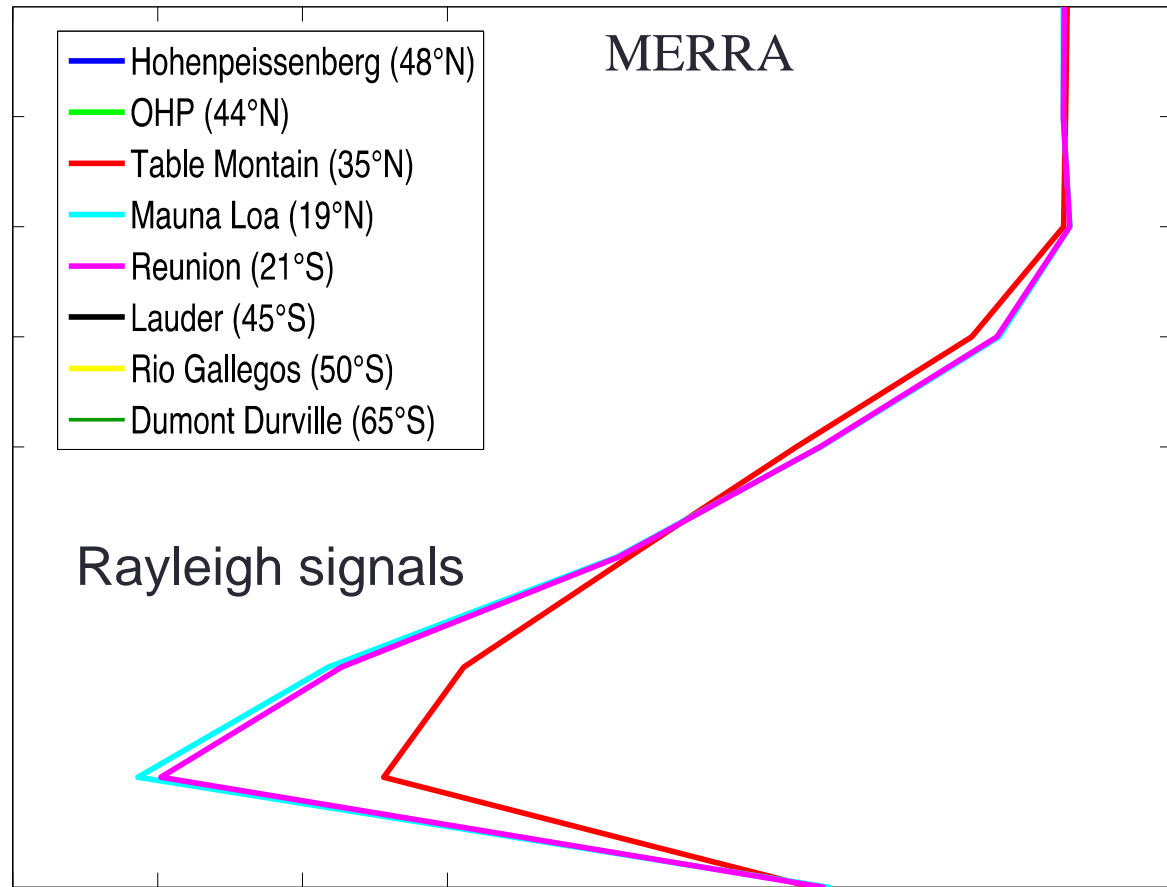
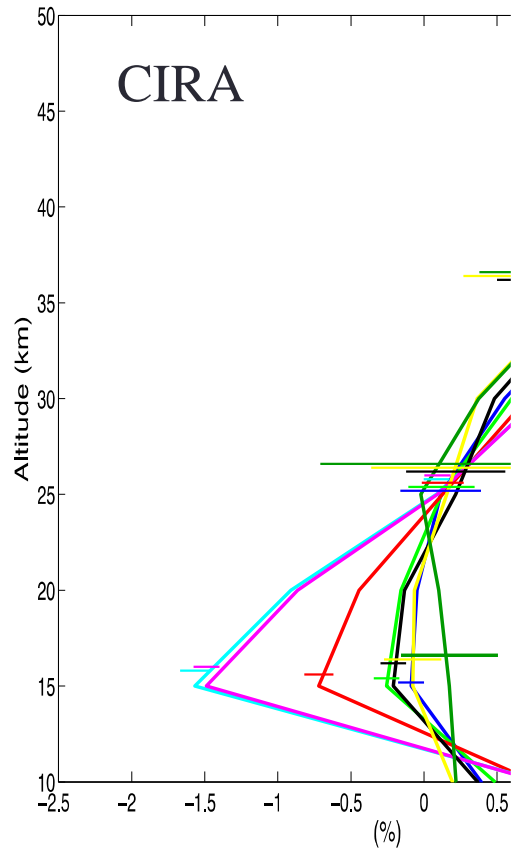


# Relative ozone difference (BP - BDM ) using different temperature models at NDACC lidar latitudes

## Temperature models

CIRA : altitude from 0 to 80 km (16 levels), spatial :  $10^\circ$  , mean zonal  
 MERRA : altitude from 1000 to 0.1 hPa (12 levels) , spatial :  $1.25^\circ \times 1.25^\circ$

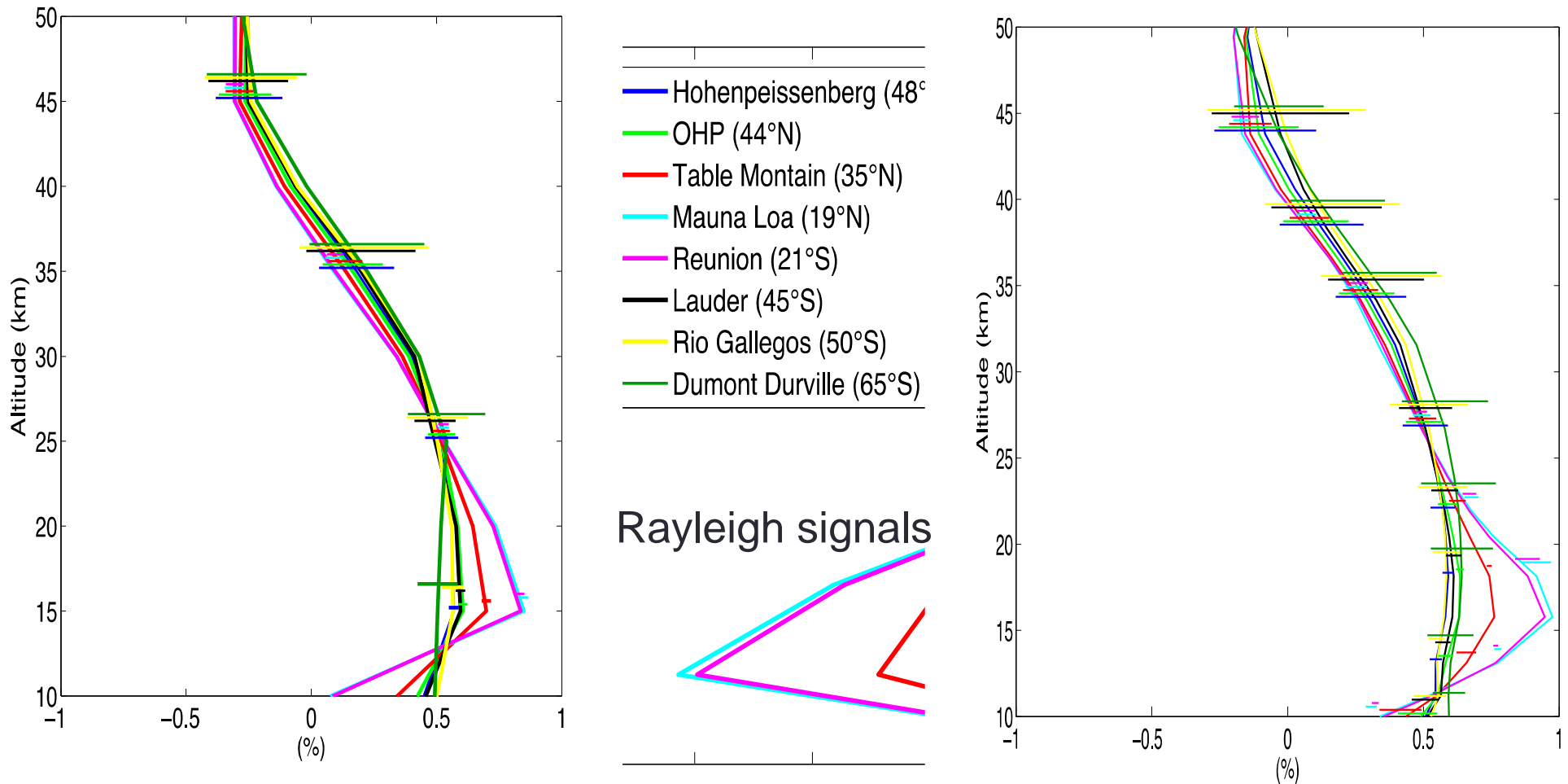
Error bars = annual variations



Rayleigh signals

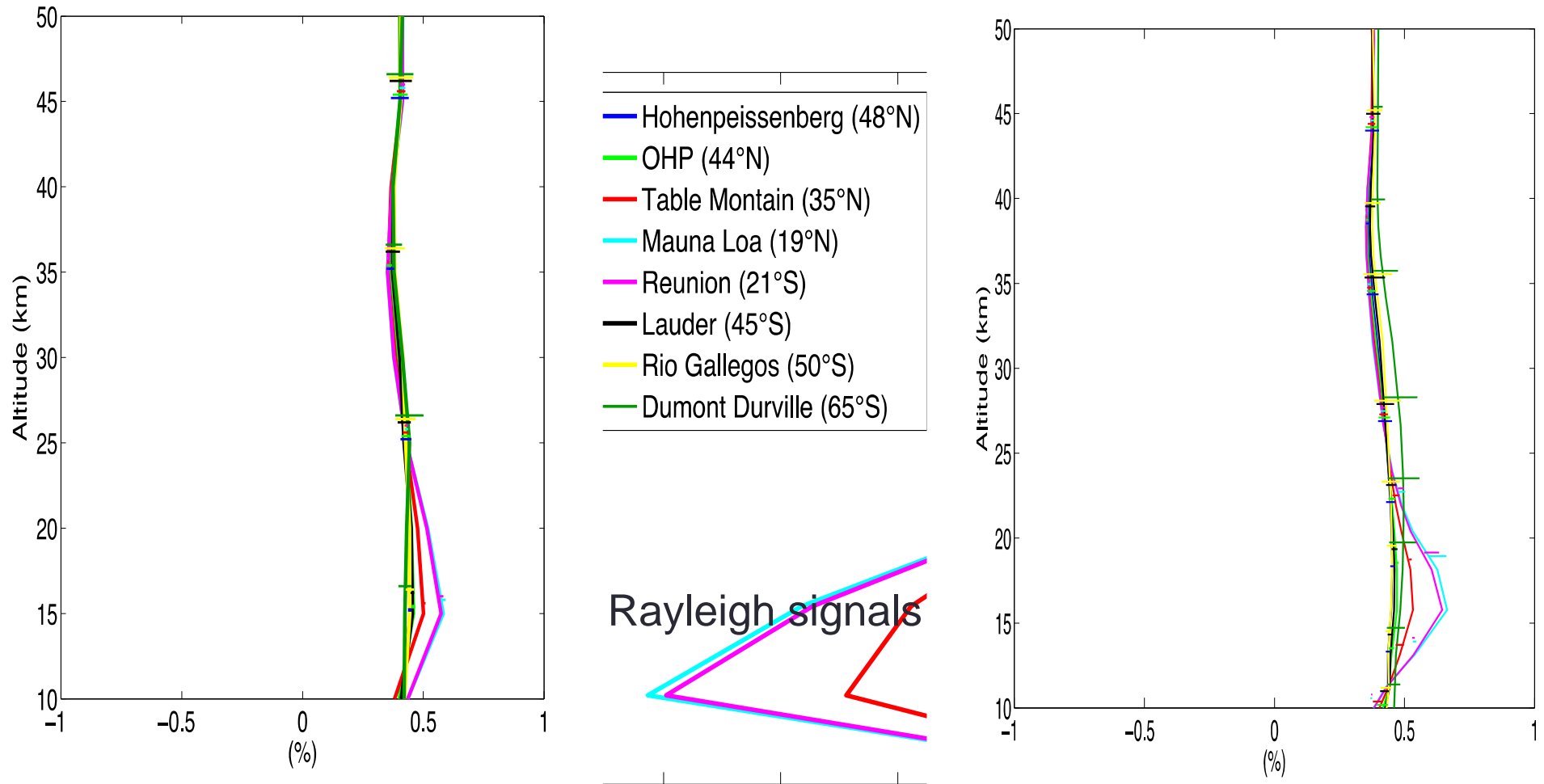
➔ In the tropics and Table M  
 Other stations : simil

# Relative ozone difference (BP – Serdyuchenko) using temperature models at various lidar NDACC latitude bands



At each stations (except in the tropics) : similar variation for both T models  
 In the tropics: max bias at ~ 15 km, more pronounced with MERRA (~ 0.9 %)

# Relative in ozone difference (BP – Average Serdyuchenko) using temperature models at various lidar NDACC latitude bands

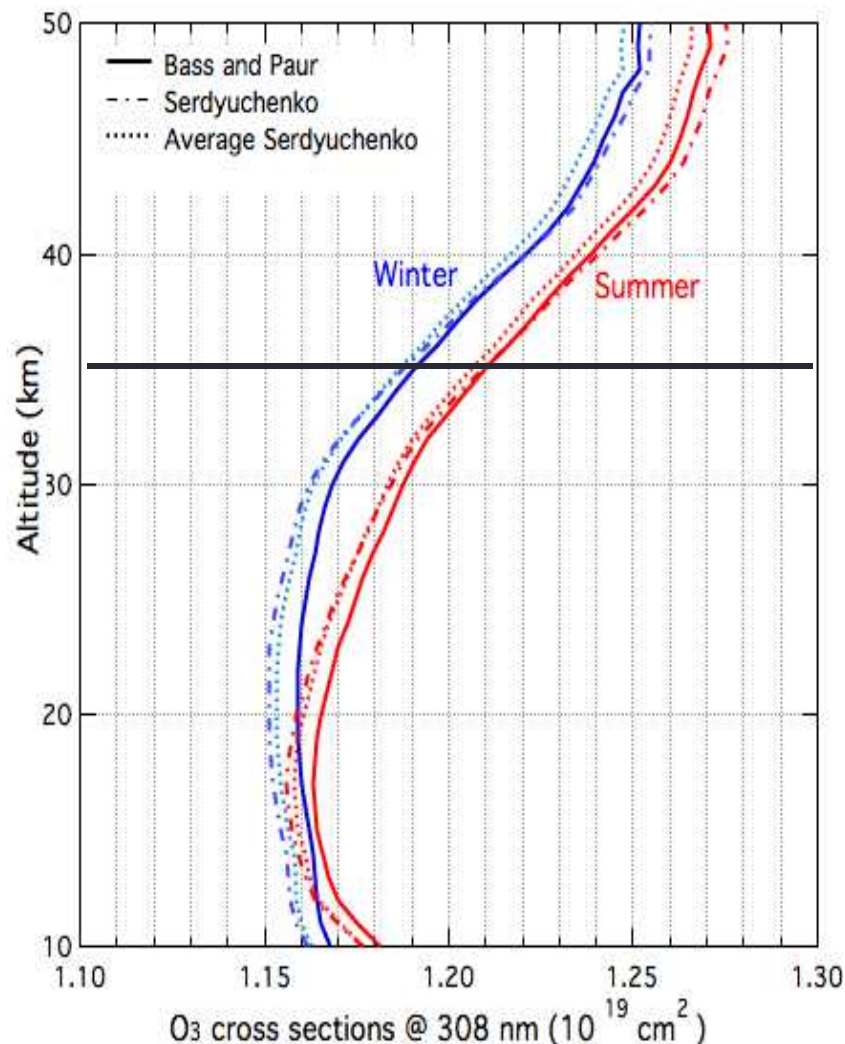


➔ At each stations (except in the tropics) : constant bias (~ 0.5 %) for both T models  
In the tropics: max bias at 15 km, more pronounced with MERRA (~ 0.6 %)

# Variation of $\sigma_{O_3}(\lambda)$ with temperature and altitude at OHP

## Temperature data used:

meteorological radiosoundings (10 to 25 km)  
climatological model: NCEP( 25 to 40 km) + CIRA (40 to 60 km)



For each cross sections : Sensitivity at 308 nm  
~ 0.2%/K

## From 10 to 35 km:

Both Serdyuchenko sections: Similar variations  
smaller than Bass and Paur

## Above 35 km:

Serdyuchenko cross sections at 308 nm: larger  
than Bass and Paur (0.3%)

Average Serdyuchenko sections: constant bias  
with Bass and Paur

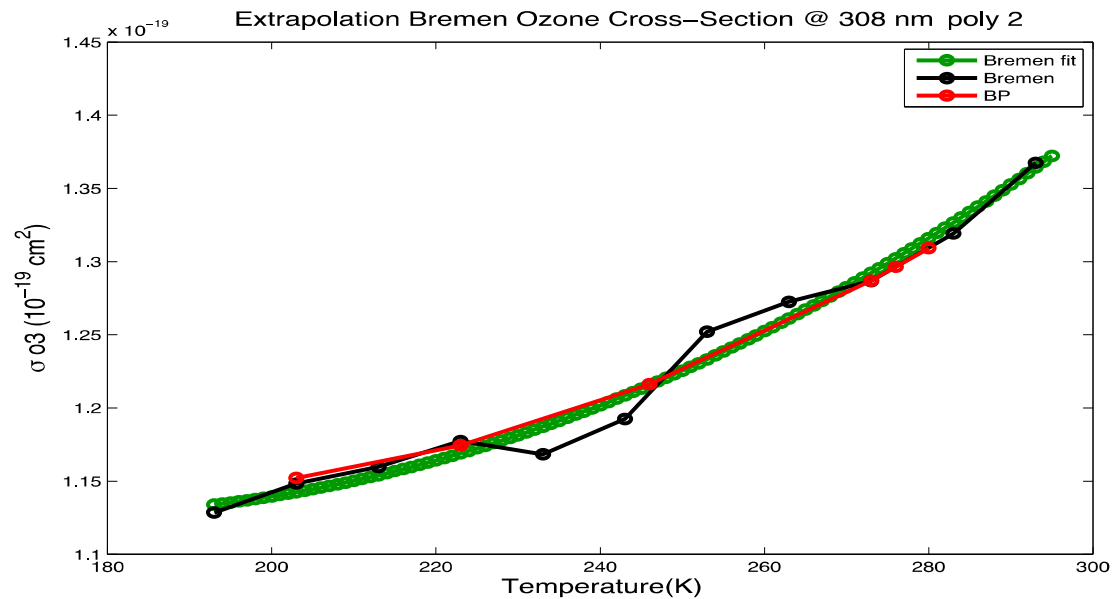
# Conclusions

- Serdyuchenko cross section at 308 nm noisier as a function of temperature than average Serdyuchenko (307.9 -308.2)
- The difference on ozone between BP and BDM in both the classical DIAL retrieval based on elastic scattering and the retrieval based on Raman scattering is below 1.5 % in absolute value from 10 to 30 km.
- The difference on ozone between BP and Serdyuchenko is below 0.5 % in absolute value and around 0.5 % for Averaging Serdyuchenko from 10 to 50 km except in the tropics

## At OHP:

- From the temperature data set used at OHP, each cross sections present an avg sensitivity at 308 nm of  $\sim 0.2\%/K$  consistent with the literature.
- Above 30 km, the difference estimated between Bass and Paur and Serdyuchenko on the  $O_3$  profiles can reach a maximum of 0.3% .
- The used of the fit average Serdyuchenko cross section can be considered for the Lidar at OHP.

# Extrapolation Bremen (308 nm)



# Extrapolation AVG bremen (307.9-308.2)

